Age of a prehistoric “Rodedian” cult site constrained by sediment and rock surface luminescence dating techniques

R. Sohbati a, b, *, A.S. Murray a, N. Porat c, M. Jain b, U. Avner d

a Nordic Laboratory for Luminescence Dating, Department of Geoscience, Aarhus University, DTU Risø Campus, Roskilde 4000, Denmark
b Center for Nuclear Technologies, Technical University of Denmark, DTU Risø Campus, Roskilde 4000, Denmark
c Geological Survey of Israel, 30 Malkhe Israel Street, Jerusalem 95501, Israel
d Dead Sea-Arava Science Center, Pato 655, Eilat 88133, Israel

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The construction age of a pavement in a “Rodedian” prehistoric cult site in Negev desert, Israel, is determined by measuring the burial age of (i) a cobble used in the pavement, and (ii) the underlying sediment. The quartz OSL age and the K-feldspar corrected IR50 age from the sediment and the corrected IR50 and pIRIR225 ages from the cobble surface are all consistent, and give an average age of 4.22 ± 0.06 ka. Although the very similar ages indicate the reliability of the methods, these ages are ~3–4 ka younger than that expected for the Rodedian sites. The IR50 and pIRIR225 luminescence-depth profiles from the cobble indicate multiple exposure and burial events in the depositional history. The apparently young ages may thus represent a later intervention in the site during the late 3rd millennium BC. More sites need to be dated by the use of both rocks and sediments to confirm this suggestion. Important information on the bleaching history of the rock surfaces directly obtained from these luminescence-depth profiles is not available in the underlying unconsolidated sediments. This is a significant advantage of rock surface dating over more conventional sediment dating.

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1. Introduction

The settlement history of the Negev desert, Israel, has been traditionally described as a series of short periods of settlement interrupted by longer gaps and archaeological sites were generally interpreted as temporary, seasonal and short-lived. However, with the advance of studies in the Negev, the settlement scenario is changing and gaps are shrinking (Avner, 2006). Of the Early Neolithic period (i.e. Pre-Pottery Neolithic periods B or PPNB, 8000–6400 BC) only about 30 habitation sites are currently known in the entire Negev (Barzilai and Goring-Morris, 2013) while almost no habitation sites need to be dated by the use of both rocks and sediments to confirm this suggestion. Important information on the bleaching history of the rock surfaces directly obtained from these luminescence-depth profiles is not available in the underlying unconsolidated sediments. This is a significant advantage of rock surface dating over more conventional sediment dating.

(Avner et al., 2014). The sites are named here “Rodedian” following the name of Nahal (Wadi) Roded in the Eilat Mountains (~5 km NW of the Gulf of Aqaba), where the highest density of cult sites was found, up to 42 sites on an area of 0.8 km2 (Fig. 1). Establishing an absolute chronology for these sites has usually relied on the 14C dating of associated organic remains; such material is not always available, and when it is, the association cannot always be assumed to be secure. Optically stimulated luminescence dating of rock surfaces has attracted significant interest in the last few years (Vafadou et al., 2007; Sohbati et al., 2011; Simms et al., 2011; Liritzis et al., 2013; Sohbati, 2015). The great advantage of this technique over the more conventional luminescence dating of buried sediments is that rock surfaces usually record their depositional history. Habermann et al. (2000), Polikreti et al. (2002), Liritzis et al. (2010) and Sohbati et al. (2011, 2012a) have all measured luminescence signal with depth into rock surfaces that were exposed to light immediately before measurement. However, by measuring such luminescence profiles into a buried rock surface, one can determine the degree to which the surface had been bleached prior to burial (Sohbati et al., 2012a,b; Chapot et al., 2012, Pederson et al., 2014) and even estimate, by using an appropriate calibration, how long the surface had been exposed before burial.
In this study we report the dating of a cult site at Naḥal Roded, using optically stimulated luminescence (OSL) of quartz, and infrared stimulated luminescence (IRSL) of potassium-rich feldspar fractions from the dust accumulated underneath a stone pavement. We support the chronology of the sediment sample by IRSL dating of whole-rock slices from the buried surface of the rock.

2. The site and sampling

The site is located on a low Precambrian igneous ridge, formed by a quartz-porphyry dike. It contains several stone features, including a pavement, on which an anthropomorphic limestone image was found lying horizontally (Fig. 2). Flint flakes collected at the site suggest Late Neolithic date (i.e. 6th millennium BC). A small paving cobble (15 × 10 × 8 cm³) and the dust accumulated underneath were collected under cover and placed in opaque plastic bags to prevent sample exposure to light. The chronological sediment sample by IRSL dating of whole-rock slices from the buried surface of the rock.

3. Sample preparation and measurement facilities

The sediment sample was wet sieved into the size fraction 63–90 μm; this fraction is probably dominated by aeolian material, and presumed most likely to have been sufficiently bleached before deposition. The grains were treated with 10% HCl to remove carbonates and 10% H₂O₂ to dissolve any reactive organic material. They were then etched with 10% HF for 40 min to remove the alpha irradiated surface layer and weathering products and coatings, followed by 10% HCl for 20 min to remove any fluoride which may have precipitated. The K-rich feldspar fraction was then separated by floating in heavy liquid (ρ = 2.58 g cm⁻³). Quartz grains (ρ > 2.58 g cm⁻³) were further cleaned and etched using 40% HF for 1 h, followed by 10% HCl for 40 min. The resulting fractions were then cut into slices of ~1.5 mm thick by a water-cooled low-speed wafering saw equipped with a 0.3 mm thick diamond blade. The surface slices (<1.5 mm depth) were treated in a similar way to the K-rich feldspar grains (using 10% HF for 40 min and 10% HCl for 20 min) to ensure the removal of any surface weathering products. No acid treatment was done on inner slices.

All luminescence measurements were carried out using Risø OSL/TL reader (Model TL-DA 15). The blue (470 nm, 80 mW·cm⁻²) stimulated signal from quartz grains was detected through a Hoya U-340 glass filter and the infrared (875 nm, 135 mW·cm⁻²) stimulated signal from K-rich feldspar grains and intact cobble slices was measured through a Schott BG39/Corning 7-59 filter combination (2 and 4 mm, respectively). Beta irradiations used a ⁵⁷⁳⁰Sr/⁹⁰Y source mounted on the reader and calibrated for both disks and cups using 180°-360° calibration quartz grains (Bøtter-Jensen et al., 2010; Hansen et al., 2015). To calibrate the source for rock slices, six slices cut from a quartzite cobble were sensitized and stabilized using successive cycles of dosing (~50 Gy) and heating (up to 450 °C). They were then given an accurately known dose of 4.81 Gy using a ¹³⁷Cs gamma source in a scatter free geometry and measured using a single-aliquot regenerative (SAR) protocol (Murray and Wintle, 2000). Slices were directly placed on the disk positions in the carousel, whereas quartz grains were mounted as large (8 mm) and K-rich feldspar grains as small (2 mm) aliquots in a monolayer using silicone oil on stainless steel disks (quartz) and cups (feldspar). The heating rate was 5 °C·s⁻¹ throughout. All thermal treatments
and stimulations at temperatures higher than 200 °C were carried out in nitrogen atmosphere. In order to allow the entire slice volume to reach the measurement temperature, the slices were held at the stimulation temperature for 30 s before the measurement started. This pause was 5 s for steel disks and cups. Five empty channels were inserted before and after the stimulation to monitor the isothermal thermoluminescence (TL).

4. Dosimetry

The radionuclide concentrations $^{238}$U, $^{226}$Ra, $^{232}$Th and $^{40}$K in both the cobble and sediment samples were measured using high-resolution gamma spectrometry. The cobble and sediment sample (dried at 50 °C) were first pulverized and homogenized. The sediment sample was heated to 450 °C for 24 h to remove any organic matter. The material was then cast in wax to prevent radon loss and to provide a reproducible counting geometry. Samples were stored for at least three weeks to allow $^{222}$Rn to reach equilibrium with its parent $^{226}$Ra before being measured on a high-purity Germanium detector for at least 24 h. Details of the gamma spectrometry calibrating are given in Murray et al. (1987). The internal beta dose rate activity from $^{40}$K in K-rich feldspar grains was calculated based on an assumed effective potassium content of 12.5 ± 0.5% (Huntley and Baril, 1997), and the beta contribution from $^{87}$Rb was calculated assuming a $^{87}$Rb content of 400 ± 100 ppm (Huntley and Hancock, 2001). A small internal alpha contribution of 0.10 ± 0.05 Gy ka$^{-1}$ from internal $^{238}$U and $^{232}$Th was also included in the dose rates, derived from $^{238}$U and $^{232}$Th concentration measurements by Mejdahl (1987). For quartz, an internal dose rate of 0.010 ± 0.002 Gy ka$^{-1}$ was assumed (Vanderbenghe et al., 2008). The radionuclide concentrations were converted to dose rate data using the conversion factors from Guérin et al. (2011). The contribution from cosmic radiation to the dose rate was calculated following Prescott and Hutton (1994), assuming an uncertainty of 5%. The long-term water content of the sediment sample was assumed to be 10% of the saturation water content. The saturated water content of the cobble is negligible. Water content, radionuclide concentrations and dry, infinite-matrix beta and gamma dose rates are summarized in Table 1.

4.1. Dose rate modelling

In order to calculate the effective dose rate at the cobble/sediment interface, the variation of beta, gamma and total dose rates with depth into both the cobble and sediment must be determined. Since the sampled cobble was part of a pavement, surrounded by other rocks of similar lithology, we model it as a homogeneous medium. The total dose rate at the rock/sediment interface of interest would have only a finite extent. Beginning by assuming the sediment empty channels were inserted before and after the stimulation to monitor the isothermal thermoluminescence (TL).

\[ D_{\beta,\text{rock}}^{\text{eff}} = 0.5D_{\beta,\text{rock}}e^{-ax} \]  

(1)

where $D_{\beta,\text{rock}}$ (Gy ka$^{-1}$) is the infinite-matrix beta dose rate of the cobble, $a$ (mm$^{-1}$) is the beta linear attenuation coefficient in sediment and $x$ (mm) is the depth into sediment from the interface. Similarly, the variation of beta dose rate in the cobble due to its own internal radioactivity is given by:

\[ D_{\beta,\text{rock}}^{\text{eff}} = D_{\beta,\text{rock}}\{1 - 0.5\left( e^{-bx} + e^{-b(h-x)} \right) \} \]  

(2)

where $b$ (mm$^{-1}$) is the beta linear attenuation coefficient in the cobble, and $h$ (mm) is the thickness of the cobble layer. Similar expressions can be used to describe the gradient of the sediment beta dose rate in sediment and rock, and the gradient of the gamma dose rates into each medium arising from both rock and sediment, using the apparent infinite-matrix dose rates and linear attenuation coefficients. We calculated the beta linear attenuation coefficient in each medium, using a weighted average of the attenuation factors for uranium and thorium series reported by Aitken (1985); Appendix H for a density of 2 g cm$^{-3}$, and an assumed attenuation factor of 1.50 mm$^{-1}$ for potassium. These attenuation factors were then modified by the assumed sediment and rock densities (2 and 2.6 g cm$^{-3}$, respectively). The gamma linear attenuation coefficients were calculated by density modification of the coefficient obtained by fitting the variation of the weighted average of fractional doses of uranium, thorium and potassium with depth given by Aitken (1985; Table H.1). The calculated gamma and beta linear attenuation coefficients for each medium are given in Table 1. The overall variation of beta and gamma dose rates was then calculated using the principle of superposition (Aitken, 1985). Fig. 3 shows the variation of beta, gamma and total dose rates through the sediment, cobble pavement and air. The total effective dose rate to surface slices from the cobble (depth <1.5 mm) and grains of sediment underlying the cobble (depth <10 mm), was calculated by integrating the total dose rate over the depth of interest.

There is no evidence that the cobble pavement has ever had a significant sediment cover. Nevertheless, had there been more than 30 cm of overlying sediment until just before sampling, the total dose rate at the rock/sediment interface of interest would have only increased by a maximum of 3%; this is considered negligible.

5. Results

5.1. Quartz

We applied a SAR protocol to measure the equivalent doses ($D_{e}$) using the OSL signal from the quartz fraction (Murray and Wintle, 2000). A preheat temperature of 260 °C for 10 s and a cut-hear (with immediate cooling) of 220 °C were used after the regeneration and test doses, respectively. All the OSL measurements were performed at 125 °C for 40 s. A further stimulation was carried out at 280 °C for 40 s at the end of each cycle to minimize the residual signal transfer between different cycles (Murray and Wintle, 2003).

Table 1

<table>
<thead>
<tr>
<th>Sample</th>
<th>Water content (%)</th>
<th>$^{238}$U (Bq kg$^{-1}$ ± SE)</th>
<th>$^{226}$Ra (Bq kg$^{-1}$ ± SE)</th>
<th>$^{232}$Th (Bq kg$^{-1}$ ± SE)</th>
<th>$^{40}$K (Bq kg$^{-1}$ ± SE)</th>
<th>Beta dose rate (Gy ka$^{-1}$ ± SE)</th>
<th>Beta attenu. factor (mm$^{-1}$)</th>
<th>Gamma dose rate (Gy ka$^{-1}$ ± SE)</th>
<th>Gamma attenu. factor (cm$^{-1}$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sediment</td>
<td>1.5 ± 1.5</td>
<td>21.8 ± 10.8</td>
<td>29.5 ± 0.9</td>
<td>34.6 ± 1.1</td>
<td>1104 ± 24</td>
<td>3.28 ± 0.08</td>
<td>1.46</td>
<td>1.48 ± 0.03</td>
<td>0.15</td>
</tr>
<tr>
<td>Cobble</td>
<td>0</td>
<td>30.2 ± 7.5</td>
<td>21.8 ± 0.6</td>
<td>44.9 ± 0.8</td>
<td>1410 ± 24</td>
<td>4.12 ± 0.07</td>
<td>1.9</td>
<td>1.79 ± 0.03</td>
<td>0.2</td>
</tr>
</tbody>
</table>
An outline of the OSL SAR protocol is given in Table 2. Signal intensities were calculated using the initial 0.8 s of the signal (channel = 1–5), less a background derived from the following 1.6 s (channel = 11–20). Data were collected in the 0.8 s preceding light stimulation (channel = 1–5) to monitor any isothermal TL signal. An early background subtraction was chosen to minimize the contribution of the more difficult to bleach and more thermally unstable medium and slow components to the net signal (Jain et al., 2003; Li and Li, 2006; Pawley et al., 2010; Cunninghan and Wallinga, 2010). We first examined the purity of our quartz extract by measuring the OSL signal from 6 aliquots with and without prior infrared stimulation at room temperature for 100 s and then calculated the ratio of the two OSL signals, the so-called infrared (IR) depletion ratio (Duller, 2003). The resulting average IR depletion ratio was 0.95 ± 0.01 (n = 6), implying that any feldspar contamination in our sample is negligible.

For the dose recovery test, six aliquots were stimulated twice at room temperature for 100 s using blue LEDs, with a pause of 1 ks between the two stimulations to allow for any charge trapped in shallow refuge traps (especially that associated with the 110 °C TL peak) to decay and subsequently partly refill the OSL trap prior to the second stimulation; the aliquots were then given doses close to their equivalent doses (D_e). The measured-to-given dose ratio was 0.94 ± 0.02 (n = 6), showing that our protocol can accurately measure (i.e. within 10%) a known laboratory dose absorbed before any thermal pretreatment. The distribution of quartz OSL D_e values is shown in Fig. 4; the average equivalent dose is 20.2 ± 1.0 Gy, which gives an age of 4.2 ± 0.4 ka (Table 3).

5.2. K-rich feldspar

K-rich feldspar is an alternative dosimeter in luminescence dating; its use is increasing following the recent identification of more stable IRSL signals than used in earlier applications (Thomsen et al., 2008). These more stable post-IR IRSL (pIRIR) signals are measured after IR stimulation at close to ambient temperature (e.g. 50 °C). In addition, Murray et al. (2012) have argued that, by comparing quartz and K-feldspar equivalent doses, it is possible to identify those samples for which the quartz OSL was well-bleached at deposition. This argument is based on the widely-observed differential bleaching rates of quartz and feldspar signals (e.g. Godfrey-Smith et al., 1988; Thomsen et al., 2008; Buylaert et al., 2012; Murray et al., 2012; Sugisaki et al., 2015; Colarossi et al., 2015) when exposed to a daylight spectrum, and helps to quantitatively address one of the main uncertainties in the application of OSL to dating.

The D_e of K-rich feldspar was measured using a SAR protocol with a preheat of 260 °C for 60 s after both regenerative and test doses. The first IR stimulation at 50 °C (IR50) was followed by a second IR stimulation at 225 °C (pIRIR225) (Buylaert et al., 2009). The pIRIR225 signal was chosen in preference to the more stable pIRIR280 as it is easier to bleach (although still substantially more difficult than IR50 or quartz OSL). A high temperature stimulation at 280 °C was also performed at the end of each SAR cycle to minimize signal carry-over to the next SAR cycle. All IR stimulations were carried out for 200 s (Buylaert et al., 2009; Sohbati et al., 2012c) (Table 2). The initial 1 s (channel = 6–10) of stimulation less a background from the last 10 s (channel = 946–995) was used for all calculations.

For dose recovery tests six aliquots were bleached using a Hönle SOL2 solar simulator for 4 h at a distance of ~80 cm (to avoid heating the samples). Three of these aliquots were then given a dose similar to the average pIRIR225 D_e value of that sample (obtained from a test run) and measured. The remaining three aliquots were also measured (without adding a dose) to determine the amount of residual dose in the samples after bleaching in the laboratory. The IR50 and pIRIR225 residual doses were 0.295 ± 0.01 (n = 6), showing that our protocol can accurately measure (i.e. within 10%) a known laboratory dose absorbed before any thermal pretreatment. The distribution of quartz OSL D_e values is shown in Fig. 4; the average equivalent dose is 20.2 ± 1.0 Gy, which gives an age of 4.2 ± 0.4 ka (Table 3).
The corrected IR50 age is 4.3 ± 0.4 ka, while the corrected pIRIR225 age of 5.9 ± 0.4 ka is apparently older (Table 4). The older pIRIR225 age is 0.3 ka which is consistent with the quartz OSL and IR50 signal and thus may not have been well bleached at the time of deposition.

5.3. Rock slices

The buried bottom surface of the cobble used in the pavement must have been concealed from light at the same time as the corresponding underlying sediment, and thus should have a burial age similar to the sediment. In order to provide an additional constraint on the age of the pavement, we also measured the dose recorded by the buried surface of the cobble.

In the absence of a practical approach to separating quartz and K-feldspar grains without losing grain-size information, we chose to work with the IR signals from solid slices taken from the surface <1.5 mm layer. Several studies have shown that the infrared stimulated blue signal from feldspars mainly originates from K-rich grains (e.g. Bail and Huntley, 2003; Tsukamoto et al., 2012; Sohbati et al., 2013), and so the same protocol used for K-rich feldspar grains was applied to solid slices (see Table 2). The shape of the IRSL signals from the slices was similar to those of K-rich feldspar grains.

To examine the applicability of our measurement protocol to whole rock slices, a dose recovery test was carried out using six surface slices (<1.5 mm) from the top surface of the cobble, exposed to daylight at the time of sampling. Three of these slices were measured to give the average residual doses of 0.3 ± 0.2 and 0.7 ± 0.3 Gy (n = 3) for the IR50 and pIRIR225 signals, respectively. The other three slices were given a dose of ~20 Gy before measurement. The average IR50 and pIRIR225 measured doses were 20.0 ± 0.1 Gy and 19.0 ± 0.1 Gy (n = 3), giving a ratio of measured doses (less residual doses) to given dose of 0.99 ± 0.01 and 0.92 ± 0.01 (n = 3), for the IR50 and pIRIR225 signals, respectively. These ratios are within 10% of unity, suggesting that our protocol for the measurement of rock slices is appropriate.

The average IR50 and pIRIR225 D0 values are 15.0 ± 0.8 Gy and 24.8 ± 0.6 Gy (n = 8), respectively (Fig. 6). Anomalous fading was measured in a similar way to K-feldspar grains. The average B2 days values for the IR50 and pIRIR225 signals were 7.5 ± 0.4 and 2.1 ± 0.2%/decade (n = 8), respectively (Table 3).

In order to provide the best estimate of the size of the luminescent K-feldspar grains in rock slices for dose rate calculation, a microscope viewfinder was calibrated using the holes (300 μm) in a single grain disc. The diameter of 40 apparently K-rich feldspar grains was then measured. 75% of these grains (mostly perthite) were larger crystals with an average size of 1200 ± 300 μm resulting in an overall average grain size of ~770 μm. This is an order of magnitude larger than the average grain size of ~77 μm (63–90 μm) for K-feldspar grains from the sediment sample. Assuming a 40K concentration of 12.5 ± 0.5% (Huntley and Baril, 2003)
and a $^{87}\text{Rb}$ content of 400 ± 100 ppm (Huntley and Hancock, 2001), the internal beta dose rate in rock and sedimentary grains is $\sim 2.8$ and $\sim 0.3\ G\text{y.ka}^{-1}$, respectively.

The IR$_{50}$ and pIRIR$_{225}$ fading corrected ages from the surface slices are 4.3 ± 0.3 ka and 4.1 ± 0.2, respectively (Table 3). These ages are indistinguishable from the quartz OSL and K-feldspar IR$_{50}$

### Table 3

<table>
<thead>
<tr>
<th>Sample</th>
<th>Quartz OSL</th>
<th>Sediment</th>
<th>Cobble</th>
</tr>
</thead>
<tbody>
<tr>
<td>$D_{e}$ (Gy)</td>
<td>$\pm$ SE</td>
<td>$\pm$ SE</td>
<td>$\pm$ SE</td>
</tr>
<tr>
<td>IR$_{50}$</td>
<td>15.0 ± 0.8</td>
<td>15.0 ± 0.8</td>
<td>15.0 ± 0.8</td>
</tr>
<tr>
<td>pIRIR$_{225}$</td>
<td>24.8 ± 0.6</td>
<td>24.8 ± 0.6</td>
<td>24.8 ± 0.6</td>
</tr>
</tbody>
</table>

Fig. 5. Distribution of K-rich feldspar IR$_{50}$ and pIRIR$_{225}$ equivalent doses.

### Table 4

Parameters used in the model and their obtained values from the fitting. Both data sets were fitted simultaneously using the global curve fitting function in OriginPro 9. The light attenuation factor $\mu$ was shared as a fixed, signal-independent parameter between the two fits. The $\tau_{i0}/D_e$ and $\tau_{i2}/D_e$ were considered as single parameters in order to reduce the number of variables in the model.

<table>
<thead>
<tr>
<th>Parameter</th>
<th>IR$_{50}$</th>
<th>pIRIR$_{225}$</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\mu$</td>
<td>0.61</td>
<td>0.61</td>
</tr>
<tr>
<td>$\tau_{i0}/D_e$</td>
<td>3.37 $\times 10^4$</td>
<td>7.98 $\times 10^4$</td>
</tr>
<tr>
<td>$\tau_{i2}/D_e$</td>
<td>706</td>
<td>1274</td>
</tr>
</tbody>
</table>

Fig. 6. Distribution of IR$_{50}$ and pIRIR$_{225}$ equivalent doses from the surface slices (<1.5 mm) from the buried bottom side of the cobble.
ages from the underlying sediment (Table 3). Considering the significant difference between the bleaching rates of these signals, the great consistency between the ages implies that the IR50 and pIRIR225 signals in the cobble surface must have been as well reset as the OSL and IR50 signals in the underlying sediment, while the pIRIR225 signal from the sediment may suffer from incomplete bleaching.

6. Was the cobble surface bleached?

In order to check whether the luminescence signal at the surface of the cobble was fully zeroed before it was incorporated into the pavement, we examined the variation of the natural sensitivity-corrected IR50 and pIRIR225 signals ($L_n/T_n$) with depth into the buried side of the cobble by measuring slices (1.5 mm thick) from three cores (of different lengths) drilled into the cobble (Fig. 7). Unfortunately, due to the shape and fragility of the cobble, we were unable to drill cores deeper than 25 mm into this sample, and the longest core broke during slicing, so that slices between 17 and 22 mm were lost. Nevertheless, the available data are very informative. Both IR50 and pIRIR225 profiles show a gentle rise, followed by flattening and subsequently a steep rise towards signal saturation with depth. However, the normalized pIRIR225 natural signals ($L_n/T_n$) increase towards saturation closer to the surface than do the IR50 signals (Fig. 7). This is consistent with earlier reports that elevated temperature pIRIR signals are more difficult to bleach than the IR50 signal (e.g. Sohbati et al., 2012c; Buylaert et al., 2012; Kars et al., 2014). On closer inspection, one can see a step in the luminescence profiles of both signals; at about ~4–5 mm for the pIRIR225 signals and ~4–10 mm for the IR50. Such steps are probably indicators of multiple burial and exposure events in the exposure history of the cobble (Freiesleben et al., 2015). The complex nature of the profiles and the variation of total dose rate with depth into the cobble (see Fig. 3), necessitates modelling for an analytical description.

Freiesleben et al. (2015) have further developed the model by Sohbati et al. (2012d) to include multiple sequential exposure and burial events in a unified equation in which the parameters have physical meaning (in contrast to simple mathematical parametrization, e.g. Laskaris and Liritzis, 2011). Following their approach of
assuming that charge trapping during exposure periods is negligible and the dose rate during discrete burial events does not change, we use the equation below to fit our data (see Fig. 8).

\[
L(x) = \left( \frac{\beta(x)}{D_0} t_{b1} + 1 \right) e^{-\tau_1 \phi \sigma e^{-\mu x}} - 1
\]

\[
L(x) = \left( \frac{\beta(x)}{D_0} t_{b2} + 1 \right) e^{-\tau_2 \phi \sigma e^{-\mu x}} - 1
\]

This equation accounts for four sequential events: two exposure events (i.e. \( t_{e1} \) and \( t_{e2} \)) each followed by a burial event (i.e. \( t_{b1} \) and \( t_{b2} \)); \( t_{e2} \) is the time elapsed since final emplacement of the cobbles in the pavement. \( \sigma \) is the luminescence decay rate at the surface of the cobbles (\( \text{ka}^{-1} \)) as a result of daylight exposure (see Sohbat et al., 2011), \( \mu \) is the light attenuation factor (\( \text{mm}^{-1} \)), and \( D(x) \) is the dose rate (\( \text{Gy, ka}^{-1} \)) which is assumed to be the same during both burial events. The parameters’ values derived from the best fit of Eqn. (3) to the data are given in Table 4.

In order to estimate the degree of bleaching of the cobbles before the final emplacement in the pavement, one can use the fitted values to predict the shapes of the profiles resulting from the last exposure event \( t_{e2} \) immediately before burial. The predicted profiles show that the IR50 and pIRIR225 signals were almost certainly completely reset to depths of ~7 and ~2 mm, respectively, before the cobbles were incorporated in the pavement (Fig. 8). Thus, we can be confident that the \( D_0 \) values measured from the surface slices do not include any poorly-bleached component (no significant residual doses) and, from that point of view at least, should be reliable.

7. Discussion

Considering the significant difference between the bleaching rates of quartz OSL and K-feldspar IR50 and pIRIR225 signals, the remarkable agreement between the corrected IR50 and pIRIR225 ages from the cobble surface and the OSL and corrected IR50 ages from the underlying sediment indicates that all these signals must have been well-bleached before the cobbles were emplaced in the pavement (Table 3). The older pIRIR225 age from the sediment probably indicates that this signal was not as well reset as the corresponding signal from the cobbles surface, perhaps, in part, because of the sampling depth of the sediment sample. The sediment underlying the cobbles was an inhomogeneous mixture of fine grain aeolian dust and coarse grain gruss produced by in-situ weathering of granitic cobbles. Thus, in order to obtain enough material in the fine grain range, the sediment was coarsely sampled up to a depth of ~1 cm using a flat trowel. It is certainly possible that the more difficult-to-bleach pIRIR225 signal was not completely reset in the grains coming from several mm below the surface. This is in contrast to the cobbles surface where complete bleaching of both IR50 and pIRIR225 signals is not only supported by the agreement between the ages obtained from the two signals, but also by an independent modelling of their luminescence-depth profiles (Fig. 8).

IR50 and pIRIR225 \( D_0 \) values from the dose response curves of the surface slices are \( 200 \pm 20 \) and \( 149 \pm 14 \text{ Gy} \) (\( n = 8 \)), respectively. Using these values and the values obtained from the fitting for \( t_{b1} / D_0 \) (Table 4), the IR50 and pIRIR225 burial ages of 3.7 \( \pm \) 0.5 and 5.0 \( \pm \) 0.3 ka for event \( t_{b1} \). These ages are of the same order of the last burial event (i.e. \( t_{b2} \)) of ~4.2 ka derived from the surface of the cobbles and the underlying sediment. The IR50 and pIRIR225 ages for the last burial event (i.e. \( t_{b2} \)) can be calculated in a similar way. The calculated pIRIR225 age is 4.2 \( \pm \) 0.2 ka which is in excellent agreement with the ages obtained directly from the surface slices and the underlying sediment, while the IR50 age of 6.7 \( \pm \) 0.1 is somewhat older. The differences presumably arise because the assumption that a single \( D_0 \) value applies to all aliquots is not strictly true; we take the burial age of ~4.2 ka derived from surface slices as the most reliable because the dose response curves were measured individually for these aliquots.

One could also obtain an exposure age from the luminescence depth profile if the decay rate at the surface \( \phi \sigma \) is known. One practical way to obtain an estimate of \( \phi \sigma \) is to use a sample of known exposure age for calibration (Sohbat et al., 2012b; Freiesleben et al., 2015). Although we did not have access to such a sample, the technique can still be used to estimate the relative length of the two exposure events. The presence of a step in our profiles means that the first exposure event \( t_{e1} \) must have been longer than the second one \( t_{e2} \), as it has reset the signal to a deeper depth (Fig. 8). In general, such step profiles are only produced if the later exposure event is of shorter or similar length to the previous event(s). Otherwise, the more recent event would completely overwrite the previous profile and thus delete all information on prior events.

8. Conclusions

We have established the construction age of a pavement in a “Rodedian” prehistoric cult site by determining the burial age of (i) a cobbles used in the pavement and (ii) the underlying sediment. Quartz OSL age and K-feldspar IR50 corrected age from the sediment and the IR50 and pIRIR225 corrected ages from the cobbles surface all converge toward an age of ~4.2 ka. This is in agreement with the age of other stone structures such as prehistoric animal traps known as ‘desert kites’ and carnivore traps in the area (Holzer et al., 2010; Porat et al., 2013).

The very similar ages obtained from both the cobbles surface and the underlying sediment indicates the reliability of the two techniques. However, these ages are ~3–4 ka younger than that
expected for the “Rodelian” sites. Although ad-hoc flint tools found
in the sites can be dated to the 6th–3rd millennium BC (Late Neolithic through Early Bronze IV), none of the 357 sites presently
surveyed yielded any pottery sherd dated to the time suggested by
the luminescence ages. This is the first attempt to date a “Rodelian”
site by means of OSL, and it is possible that these apparently young
OSL ages represent a later intervention in the site during the late
3rd millennium B.C. This suggestion is consistent with the exposure
and reburial events identified in the luminescence profile from the
cobble. More sites need to be dated by the use of both rocks and
sediments to confirm this suggestion.

This study adds to the increasing evidence that OSL dating of
rock surfaces is a robust and reliable dating technique. The
invaluable information on the bleaching history of the rock surfaces
that is directly obtained from luminescence-depth profiles is not
available in unconsolidated sediments. This is a significant advan-
tage of rock surface dating over more conventional sediment
dating.

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